

The influence of degree of peat decomposition on phosphorus binding forms in fens

S. Jordan, S. Velty and J. Zeitz

Faculty of Agriculture and Horticulture, Institute of Crop Science, Humboldt-Universität zu Berlin, Germany.

This article is dedicated to the memory of our very good friend and colleague,

Dr Silke Velty, 07 April 1973 – 14 October 2007.

SUMMARY

Re-wetting of drained fens can release phosphorus, introducing a eutrophication risk for associated aquatic ecosystems. Characterisation of the different forms of organic and inorganic bound phosphorus in the peat is an important step towards the development of tools for assessing the level of risk attached to individual re-wetting projects. In the work reported here, a sequential extraction (fractionation) method was used to distinguish the following P binding forms:

1. labile P, detected by NH_4Cl extraction;
2. redox-sensitive P, detected by $\text{Na}_2\text{S}_2\text{O}_4/\text{NaHCO}_3$ extraction;
3. P adsorbed to metal oxides, detected by HCl extraction;
4. P bound to humic substances, detected by NaOH extraction; and
5. organic and refractory bound P, detected using H_2SO_4 and H_2O_2 .

Special attention was paid to the degree of decomposition (DPD) of the peat, and metal concentrations were measured in selected fractions. Higher P concentrations were found in completely humified than in little humified peat for all fractions except the NH_4Cl (labile P) fraction, where P content increased as DPD decreased. As only 1% of total phosphorus (TP) was present as labile P, the results indicate that the decisive horizons for nutrient release after re-wetting are those that are completely humified due to pedogenetic changes. The principal metal sorption partner for P was Fe.

KEY WORDS: fen restoration, humification, peat soil, re-wetting, sequential extraction.

INTRODUCTION

In north-eastern Germany, the drainage of huge areas of fen (*sensu* Joosten & Clarke 2002) over several decades means that their natural functions as accumulators, buffers, transformers and filters for nutrients and water have been impaired. Fundamental changes in the hydrology, chemistry, biochemistry, physics and biology of the peat are apparent. Aerobic conditions have resulted in oxidative peat decomposition, with concomitant emission of CO_2 to the atmosphere. Possible consequences are peat mineralisation and transformation of organically bound phosphorus (P) into redox-sensitive inorganic Fe(III) oxyhydroxide-bound P (Gelbrecht & Koppisch 2001, Zak *et al.* 2004). The porosity, water capacity, hydraulic conductivity and wettability of the peat have all been reduced (Zeitz 2003) and the peat body is characterised by structural damage (Lehrkamp 1989, Succow & Jeschke 1990, Velty *et al.* 2006, Zeitz & Velty 2002) such as aggregation in the peat shrinkage horizon (Figure 1) with adherent shrinkage cracks (Figure 2).

The regeneration of fens is now promoted both economically and politically, for example by the Kyoto Protocol (UNFCCC 1998) and the European Water Framework Directive (EU 2000). Many degraded fens have been re-wetted and there are plans to re-wet more in the near future. Re-wetting leads to saturated conditions which prevent peat decomposition and may initiate new peat growth.

Re-wetting of degraded fens is often accompanied by enhanced release of P due to changes in the redox potential of P-adherent chemical elements such as Fe (Paludan 1995, Gelbrecht & Koppisch 2001, Gelbrecht *et al.* 2003). Thus, in managing degraded fens, special account must be taken of the potential for supplementation of nutrient outputs to adjacent aquatic ecosystems. Partitioning, mineralisation, immobilisation, transport and uptake of P are well-known individual processes that govern P dynamics and whose effects on surface water eutrophication have been studied worldwide (Bar-Yosef 2003).

Total phosphorus (TP) concentrations in natural and uninfluenced waters are mostly below 0.1 mg l^{-1} (Rump 1998). In re-wetted fen sites Rupp *et al.*



Figure 1. Peat aggregation (crumbs) in the peat shrinkage horizon of a fen.



Figure 2. Shrinkage cracks in fen peat due to aerobic conditions associated with agricultural use.

(2004) report soluble reactive P (SRP) concentrations up to 0.36 mg l^{-1} in surface water whilst Gelbrecht & Zak (2004) measured SRP concentrations up to 18.9 mg l^{-1} in pore water.

In order to obtain more detailed information on the P pool that could be released after re-wetting of fens, it is important to assess accumulated P in the peat. Sequential extraction of P from peat soil samples provides information about its potential mobility under different environmental conditions; and the determination of P binding forms helps to clarify mechanisms of P release likely to result from re-wetting so that P release potentials may be estimated prospectively.

Conventional investigation methods such as the Calcium-Acetate-Lactate-Digestion or the Double-Lactate-Digestion (VDLUFA 1991, Scheffer & Schachtschabel 2002) are used for the determination of plant-available P. The accuracy of these methods is sufficient for the evaluation of agricultural fertiliser requirements, but no information about soil-chemical complexes is obtained because the recorded P fraction cannot be related plausibly to the binding forms. The first sequential extraction methods for soils were developed e.g. by Dean (1938), Chang & Jackson (1957), Smith (1965) and

Hedley *et al.* (1982) and these procedures have been adapted for the analysis of P in lake sediments e.g. by Hieltjes & Lijklema (1980), Pettersson (1986) and Psenner *et al.* (1988). In the sequential extraction after Chang & Jackson (1957) which has been applied, modified and simplified worldwide (Tiessen & Moir 1993, Guppy *et al.* 2000, Tyler 2002), organic P is detected only as the difference between TP and the sum of all inorganic P forms (Pagel *et al.* 1982, Barbanti *et al.* 1994).

For fens, there is no specific sequential extraction method that allows differentiation of P binding forms, and the processes of P shifting and P release are still incompletely understood. This situation creates difficulties in estimating the risk of re-wetting for the environment. More studies on the influence of the degree of peat decomposition (DPD) on P release are also needed before a general risk assessment scheme for peat substrates can be devised.

In the work described here, we tested the hypothesis that DPD influences the P binding forms in fens. We determined inorganic and organic P binding forms by sequential extraction and related this to DPD. We also determined the content of metals as potential sorption partners for P.

METHODS

Peat and degree of decomposition

Representative peat samples (n=16) with different DPD were taken from a fen in the Peene Valley (Mecklenburg-Western Pomerania, Germany). Sample selection was based on von Post's humification scale, which involves visual estimation of the colour and turbidity of the water that can be extracted by squeezing the peat through the fingers. As a rule, the clearer (without residuum) the water, the less humified is the peat (von Post 1924, von Post & Granlund 1926, Hämäläinen 1991). Degree of decomposition (humification) is expressed in terms of a ten-class scale on which higher numbers indicate stronger peat decomposition (Table 1).

Table 1. Degree of peat decomposition after von Post, and number of peat samples analysed.

Degree of decomposition after von Post	Level of humification	Number of peat samples
H 1	completely/virtually unhumified	
H 2		
H 3	little/poorly humified	6
H 4		2
H 5	fairly humified	2
H 6		
H 7	quite well-/well humified	2
H 8		
H 9	almost completely/completely humified	
H 10		4

Whilst von Post's method for estimation of DPD is a subjective means for rapid field determination

of the degree of humification, it is indispensable in this context and should complement rather than be displaced by "modern" procedures.

Sequential extraction of phosphorus

Knowledge of the P forms in soils has been derived largely by chemical fractionation based on the ability of selective chemical reagents to dissolve specific types of organic and inorganic compounds.

For the present study, the sequential extraction method of Psenner *et al.* (1984), originally used for the detection of P binding forms in lake sediments (Jensen & Thamdrup 1993, Hupfer 1995, Paludan 1995, Lewandowski 2002), was adapted to take account of the special properties of peat.

The fractionation was carried out on replicate field-fresh peat samples as soon as possible after collection, since drying induces oxidation of the peat and leads to irreversible changes in its fractional composition.

A 5 g sub-sample of each peat sample was placed in a 25 ml centrifuge tube and extracted sequentially into five fractions. At each stage, bound P was assumed to be removed selectively from specific types of compounds contained in the peat (Table 2 and Figure 3).

The bio-availability of the prevalent P binding forms decreases with each extraction step (Pettersson 1986). For this reason, stable organic and refractory bound P was determined as "Rest-P" (residue of the NaOH fraction) in the final extraction step, which involved washing with deionised water, decanting, drying, grinding and taking a 20 mg sample for H₂SO₄/H₂O₂ digestion.

In order to verify the P content of the fractions, it was necessary also to determine TP for each peat sample by H₂SO₄/H₂O₂ digestion (10 mg of dry peat) (Zwirnmann *et al.* 1999, DIN EN 1189

Table 2. Fractionation scheme showing the five extraction steps, the extraction solvent and P binding form (BD = Bidithionite, SRP = soluble reactive phosphorus, NRP = non-reactive phosphorus, TP = total phosphorus).

Fraction	Extraction solvent	P binding form
NH ₄ Cl-SRP	1 M NH ₄ Cl	labile, loosely bound or adsorbed P; immediately available
BD-TP	0.1 M NaHCO ₃ -Na ₂ S ₂ O ₄ -Solution	redox-sensitive P, mainly from Fe(III)- and Mn(IV)-oxyhydroxide surfaces
HCl-SRP	0.5 M HCl	P adsorbed to metal oxides (Al-, Fe- and Mn-bound P); apatite-P and P bound in carbonates
HCl-NRP		acid labile organic P
NaOH-TP	1 M NaOH	humic substance bound P and poly-P
Rest-P	TP detection	organic and refractory bound P

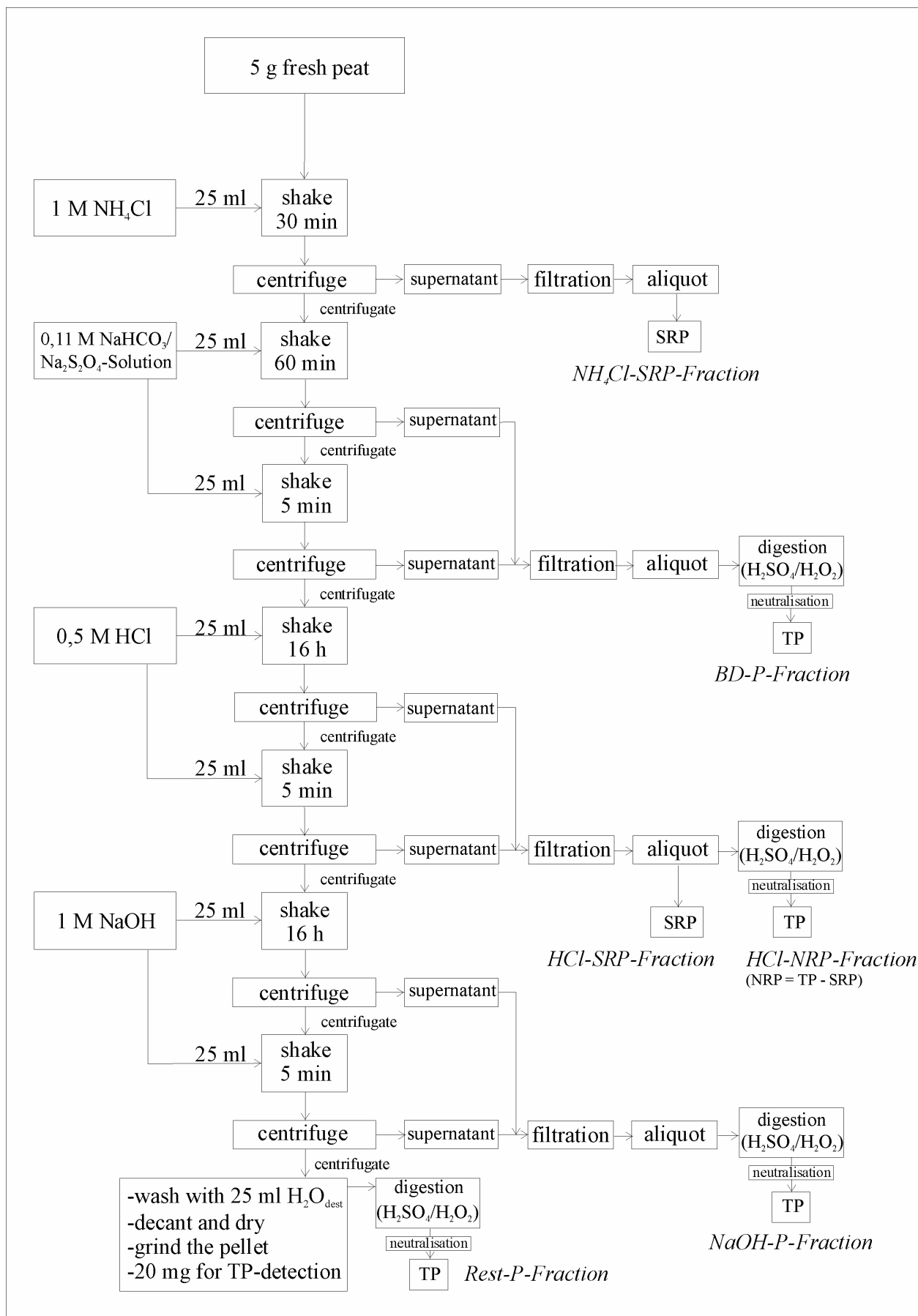


Figure 3. Procedure of the modified PSENNER-Fractionation (centrifuge 5 min at 10000·g; filtration with 0.45 μm CA-Filter, Whatman).

1996). Gasparatos & Haidouti (2001) compared different methods for the detection of TP in soils and found no significant differences between the concentrations of TP extracted by digestion in HClO_4 , aqua regia and $\text{H}_2\text{SO}_4/\text{H}_2\text{O}_2$.

P concentrations in the extracts and TP were detected photometrically (at $\lambda_{\text{max}}=680$ nm) using the Molybdenum Blue Method (Murphy & Riley 1962).

The concentrations of labile P, redox-sensitive P, P adsorbed to metal oxides and P bound to humic substances are presented in this paper. Rest-P and TP are not discussed because these fractions contain stable P that is not immediately available.

Analysis of metals

Fe, Al, Mn and Ca were determined in selected fractions using an Atomic Absorption Spectrometer (Perkin-Elmer 3300 with flame). Ca was detected mainly as soluble calcium(hydrogen)carbonate within the NH_4Cl fraction; Fe and Mn were determined in the BD fraction; and Al, Ca, Fe and Mn in the HCl-SRP-fraction (Table 3). No metals were measured in the NaOH fraction or in Rest-P because the P here was bound organically and therefore refractory. Total metal concentrations for each sample were also determined.

Table 3. Scheme of extractions for metal analyses.

Fraction	P binding form	Detected metals
NH_4Cl-SRP	labile, loosely bound or adsorbed P; immediately available	Ca
BD-TP	redox-sensitive P, mainly from Fe(III) and Mn(IV) oxyhydroxide surfaces	Fe, Mn
HCl-SRP	P adsorbed to metal oxides (Al-, Fe- and Mn-bound P); apatite-P and P bound in carbonates	Al, Ca, Fe, Mn

RESULTS

The relationships between the P concentration in each fraction and DPD are shown in Figure 4. Redox-sensitive P data are presented for all 16 peat samples (Figure 4b), but only 12 data points are shown for the other P binding forms due to methodological changes. Regression analysis (calculation of R^2) was not carried out because the class variable DPD is discontinuous.

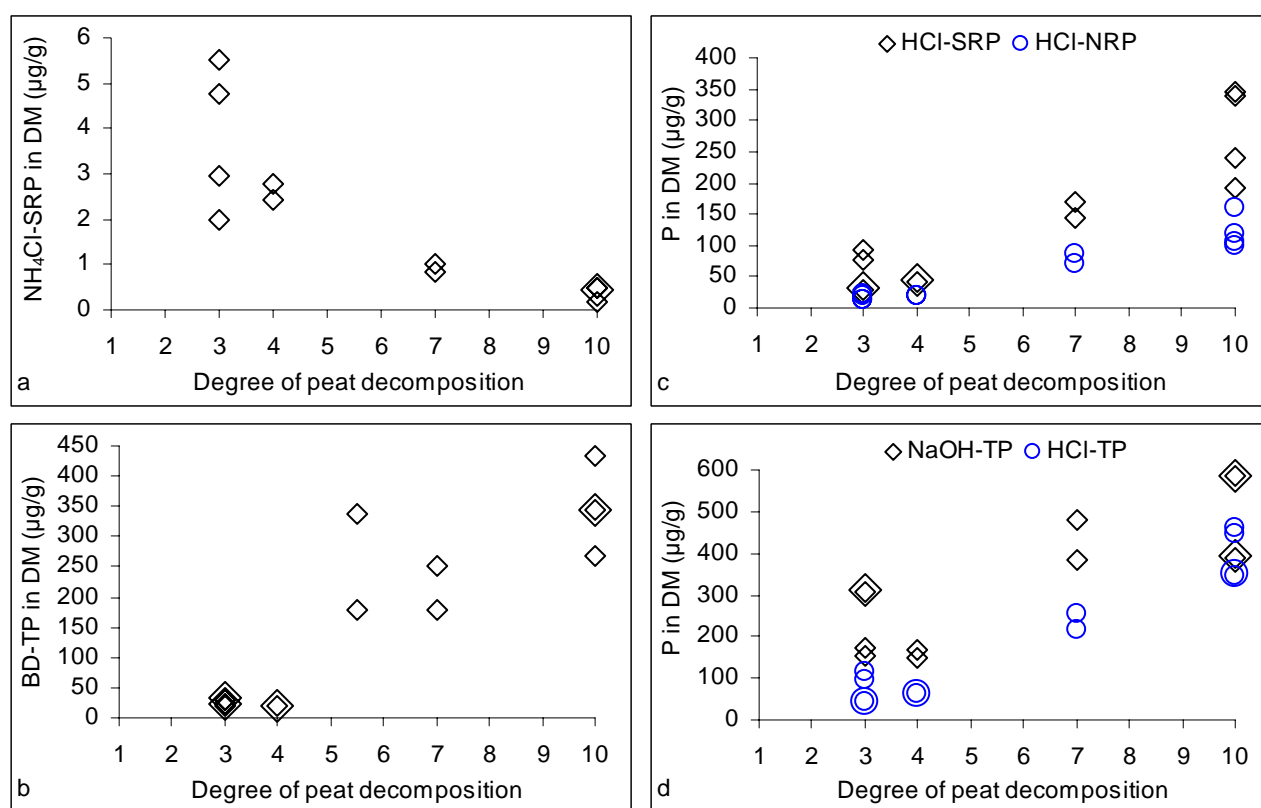


Figure 4. Relationships between P binding form concentrations and degree of peat decomposition in fen peat samples: (a) labile P (NH_4Cl -SRP); (b) redox-sensitive P (BD-TP); (c) metal oxide bound P (HCl-SRP) and acid labile organic P (HCl-NRP); (d) humic substance bound P (NaOH-TP) and HCl-TP. DM=dry matter.

Phosphorus fractionation

The concentration of labile P was lowest (NH_4Cl -SRP below $0.5 \mu\text{g g DM}^{-1}$) in the completely humified (H 10) peat horizon, and increased with decreasing DPD (Figure 4a). The highest concentration ($5.5 \mu\text{g P g DM}^{-1}$) was found in a little humified (H 3) sample.

The content of redox-sensitive P was higher than that of labile P and increased with increasing DPD, ranging from $20.7 \mu\text{g P g DM}^{-1}$ in little humified (H 3) peat to $434 \mu\text{g P g DM}^{-1}$ in completely humified (H 10) peat (Figure 4b).

The largest concentrations of metal oxide bound P (HCl -SRP: $347 \mu\text{g P g DM}^{-1}$; H 3: $27.7 \mu\text{g P g DM}^{-1}$) and acid labile organic P (HCl -NRP: $159.2 \mu\text{g P g DM}^{-1}$; H 3: $11.2 \mu\text{g P g DM}^{-1}$) were found in the completely humified (H 10) peat sample. The content of P again increased with increasing DPD in

this fraction (Figure 4c).

The concentration of humic substance bound P (NaOH -TP) increased from $152 \mu\text{g P g DM}^{-1}$ in little humified (H 3) peat up to $587 \mu\text{g P g DM}^{-1}$ in completely humified (H 10) peat (Figure 4d).

Metals

No relationships were found between DPD and Al:P , Ca:P or Mn:P . However, a statistically significant linear relationship ($R^2=0.85$) was found between the concentrations of redox-sensitive P and Fe, the P content increasing with increasing Fe content (Figure 5a). The Fe content of the BD fraction increased with DPD, from $0.64 \mu\text{mol Fe g DM}^{-1}$ at H 3 to $116.6 \mu\text{mol Fe g DM}^{-1}$ at H 10 (Figure 5b); and the Fe content of the HCl fraction increased from $6.3 \mu\text{mol Fe g DM}^{-1}$ at H 3 to $246.1 \mu\text{mol Fe g DM}^{-1}$ at H 10 (Figure 6).

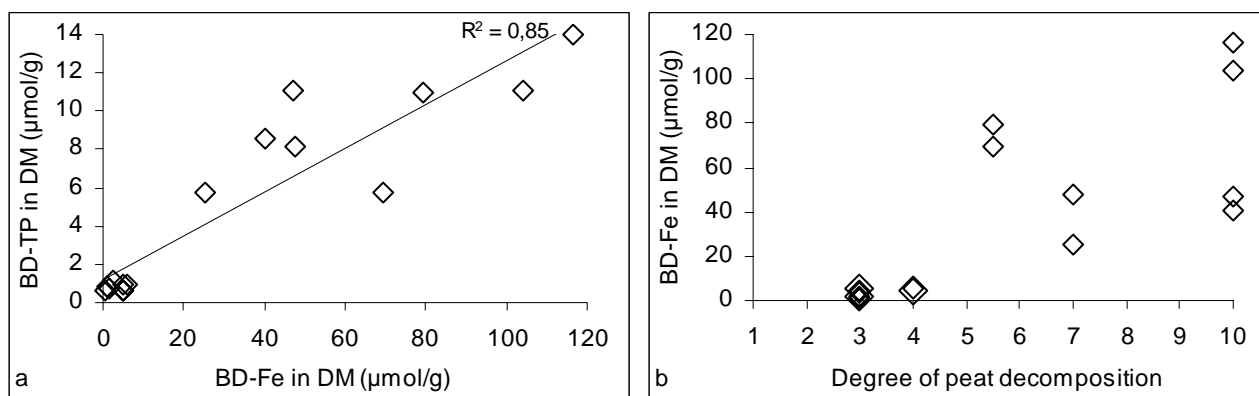


Figure 5. (a) Correlation between Fe and P in the BD fraction; and (b) relationship between Fe in the BD fraction and DPD ($n=16$).

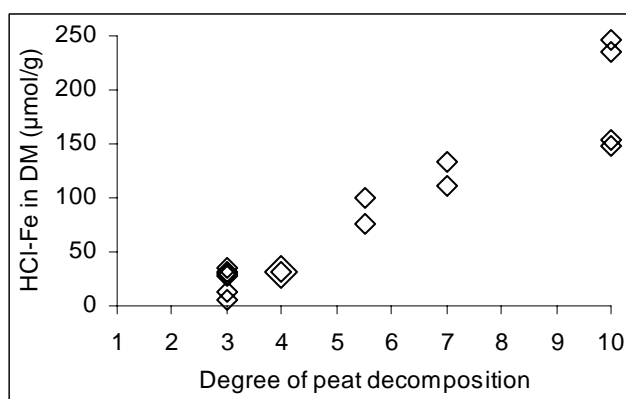


Figure 6. Relationship between Fe in the HCl fraction and DPD ($n=16$).

DISCUSSION

Applicability of methods

Sequential extraction methods were developed to better quantify the (clustered) bio-availability of different pollutants and trace elements associated with soils and sediments in a rapid and simplified way. They generally employ strong solvents and reactions, in contrast to the weak solvents and slow reactions of the natural processes they simulate. As a result, they cannot be used to identify discrete compounds because of deviations from natural physicochemical conditions and have been described as “operational” due to e.g. re-adsorptions, relocations (Lewandowski 2002) and overlaps between the fractions (Martin *et al.* 1987,

Pettersson *et al.* 1988, Psenner *et al.* 1988). Nonetheless, each extraction solvent should destroy one fraction without affecting the others, and should recover all of the theoretically assumed pool of elements. Unfortunately, expected and measured results differed dramatically in our study, with recoveries both up to 20% below and above 100% being obtained for a few of the peat samples. This may be due to the heterogeneity of the peat.

Despite these problems, sequential extraction methods remain useful for characterising interactive processes between elements in ecosystems, and they provide useful data for modelling the risks resulting from pollution.

Labile phosphorus

The fraction of TP that is present as labile P is very low (0.016 % at H 10, 1.25 % at H 3), and the highest labile P value in the NH_4Cl fraction is exceeded many times over by the P concentrations in other fractions. Therefore, P release from the immediately available labile fraction may be disregarded as a cause of nutrient release into adjacent ecosystems after re-wetting of fens. However, other chemical consequences of re-wetting are important.

Peat water directly influences P detection only in the first extraction step (NH_4Cl fraction). Thus, the large labile P content of little humified (H 3) peat might be attributable to peat water. Parent & Ilnicki (2003) report a relationship between DPD and bulk density in which peat with low degree of decomposition has the lowest bulk density. Low bulk density means, in turn, high porosity and consequently high accumulation of peat water and labile P. This could explain the reverse P distribution in the NH_4Cl fraction as compared to the other fractions.

Redox-sensitive phosphorus

A large fraction of TP is present as redox-sensitive P, especially in completely humified peat horizons (4.7 % at H 3, 34.8 % at H 10) due to the high Fe concentration in this peat horizon (see Figure 4b). The relationship is governed by the strong binding of inorganic P to Fe(III)- and Mn(IV)-oxyhydroxides. The reductive dissolution of these compounds is likely to be a significant mechanism for P release into adjacent ecosystems when degraded fens are re-wetted, and can be simulated on the basis of the data for this fraction.

Phosphorus adsorbed to metal oxides and acid-labile organic phosphorus

The concentration of HCl-SRP was higher than the concentration of HCl-NRP in all the peat samples

examined, but the relative amounts of HCl-SRP and HCl-NRP varied (Figure 4c). The ratio of labile organic P (HCl-NRP) to HCl-SRP was higher in little humified (H 3, H 4) than in completely humified (H 10) peat due to the low concentrations of P binding partners such as Fe, Mn, Al and Ca at H 3. In completely humified horizons, a large fraction of TP was adsorbed to metal oxides and acid labile organic P (HCl-TP) (33.4 % at H 10 compared with 9.5 % at H 3), and this could be released by re-wetting.

Phosphorus bound to humic substances

Organically bound P comprises 25-65 % of TP in mineral soils and 80 % in peat soils (Harrison 1987, Stevenson 1986, Scheffer & Schachtschabel 2002). The principal forms of organic P are inositol phosphates, phospholipids and nucleic acids or their degradation products, whilst sugar phosphates and phosphoproteins occur as organic P trace compounds (Anderson 1975, Harrison 1987, Stevenson 1986).

Peat with high humification has the largest content of P bound to humic substances, which is not immediately available (NaOH-P 42.7 % at H 10, 33.3 % at H 3).

Larger P contents are found in the surface horizons of fens because they are more humified and enriched with metal oxides (HCl-P) and humic substances (NaOH-P) (Figure 4d). This relationship can be explained by the fact that humic substances are formed gradually during humification, and so are present at lower concentrations in little humified than in completely humified peat (Koppisch 2001). Due to the concomitant mineralisation, organic P from dead plant material in the process of decomposition and P from fertilisers is permanently fixed to these newly-formed humic substances, to organic decomposition products, and to metal oxides. At the same time, the peat is enriched by P due to humification.

Metals

The BD fraction simulates the reduction of Fe(III) to Fe(II) with concomitant release of P during re-wetting of fens. The duration of P release is unclear, but may persist for many years. We found high BD-TP contents in peat soil samples nine years after re-wetting, with the implication that large quantities of Fe were still available in oxidised form after this period of time.

Mn(IV) oxides can be neglected as P binding partners due to the occurrence of only minor concentrations in the BD extract. Thus the results indicate that the predominant P binding mechanism involves Fe.

CONCLUSIONS

1. The results of sequential extraction indicate that only labile P, redox-sensitive P and part of the metal oxide-bound P could be released from the fen peat substrate by re-wetting.
2. Only 1% of the total phosphorus (TP) is present in labile form, and so presents no risk for increased nutrient release into adjacent ecosystems after re-wetting of fens.
3. High potential for release of phosphorus may exist in well and completely humified fen peat horizons with high Fe(III)-oxyhydroxide contents. If oxidised Fe remains after re-wetting, P may be released from these horizons over many years.
4. Phosphorus that is bound to humic substances and Rest-P can be disregarded as sources of nutrient release after re-wetting because these P binding forms are stable, organically bound and refractory, and thus not immediately available.
5. When re-wetting a drained fen, well humified horizons have to be taken into account because they have been changed pedogenetically.
6. The results of the P fractionation can be interpreted in terms of a qualified correlation with parallel detection of metals in the same extraction solvent (Hupfer 1995). Fe, Al, Mn and Ca are supposed to be potential sorption partners for P. As a rule, the content of P-binding metals in mineral soils and in peat substrates is higher than the content of P, so that all P may be bound to these elements (Tischner 2000). The solution properties of the sorption partners in the different extraction solvents differ sufficiently from one another to provide an indication of P release potential on re-wetting. No relationships of P with Al, Ca, Mn or DPD were found, so that P binding predominantly to Fe is indicated for peat samples from the locality investigated.

ACKNOWLEDGEMENTS

We thank Jörg Gelbrecht, Dominik Zak and the staff of the Central Chemical Laboratory, Leibniz Institute for Fresh Water Ecology and Fishery (IGB) for support in carrying out the chemical analyses and for interesting discussions. Louise Dunn made improvements to the text. The research was funded by the scholarship programme of the Deutsche Bundesstiftung Umwelt (DBU). This paper is based on a presentation by Sabine Jordan at the Fifth European Conference on Ecological Restoration held at Greifswald, Germany, in August 2006.

REFERENCES

- Anderson, G. (1975) Other organic phosphorus compounds. In: Gieseking J.E. (ed.) *Soil Components. Vol. 1, Organic Components*. Springer-Verlag, Berlin, 305–331.
- Barbanti, A., Bergamini, M.C., Frascari, F., Miserocchi, S. & Rosso, G. (1994) Critical aspects of sedimentary phosphorus chemical fractionation. *Journal of Environmental Quality*, 23, 1093–1102.
- Bar-Yosef, B. (2003) Phosphorus dynamics. In: Benbi, D.K. & Nieder, R. (eds.) *Handbook of Processes and Modeling in the Soil-Plant System*. Food Products Press and The Harworth Reference Press, Imprints of the Haworth Press Inc., Binghamton, NY, 483–523.
- Chang, S.C. & Jackson, M.L. (1957) Fractionation of soil phosphorus. *Soil Science*, 84, 133–144.
- Dean, L.A. (1938) An attempted fractionation of the soil phosphorus. *Journal of Agricultural Science*, 28, 234–246.
- DIN EN 1189 (1996) Bestimmung von Phosphor (Determination of phosphorus). In: Fachgruppe Wasserchemie in der Gesellschaft Deutscher Chemiker in Gemeinschaft mit dem Normenausschuß Wasserwesen im DIN Deutsches Institut für Normung e.V. (eds.) *Deutsche Einheitsverfahren zur Wasser-, Abwasser- und Schlammuntersuchung*, DEV D 11, 38, Lieferung 1997, Weinheim, 25 pp. (in German).
- EU (2000) Directive 2000/60/EC of the European Parliament and of the Council of 23 October 2000 establishing a framework for Community action in the field of water policy. Online at: http://ec.europa.eu/environment/water/water-framework/index_en.html.
- Gasparatos, D. & Haidouti, C. (2001) A comparison of wet oxidation methods for determination of total phosphorus in soils. *Journal of Plant Nutrition and Soil Science*, 164, 435–439.
- Gelbrecht, J. & Koppisch, D. (2001) Phosphor-Umsetzungsprozesse (Chemical reactions of phosphorus). In: Succow, M. & Joosten, H. (eds.) *Landschaftsökologische Moorkunde (Landscape Ecology of Peatlands)*, Second Edition, Schweizerbart, Stuttgart, 24–26 (in German).
- Gelbrecht, J., Lengsfeld, H. & Zak, D. (2003) Stoffrückhalt und -freisetzung in grundwasser- gespeisten Mooren des nordostdeutschen Tieflandes (Retention and release of substances in minerotrophic peat soils in the north-eastern lowland of Germany). In: Lenschow, U. (ed.) *Stoffausträge aus wiedervernässten Niedermooren (Discharge of substances from re-wetted*

- minerotrophic peat soils*). Schriftenreihe des Landesamtes für Umwelt, Naturschutz und Geologie (LUNG), Mecklenburg-Vorpommern, 1/2003, 55–66 (in German). http://www.lung.mv-regierung.de/dateien/stoffaustrag_v4.pdf.
- Gelbrecht, J. & Zak, D. (2004) Stoffumsetzungsprozesse in Niedermooren und ihr Einfluß auf angrenzende Oberflächengewässer (Chemical reactions in minerotrophic peat soils and their influence on adjacent surface waters). *Wasserwirtschaft*, 94(5), 15–18 (in German).
- Guppy, C.N., Menzies, N.W., Moody, P.W., Compton, B.L. & Blamey, F.P.C. (2000) A simplified sequential phosphorus fractionation method. *Communication in Soil Science and Plant Analysis*, 31, 1981–1991.
- Hämäläinen, M. (1991) *Principal Variations in the Chemical Composition of Peat*. Swedish University of Agricultural Sciences, Department of Chemistry, Uppsala, 21–22.
- Harrison, A.F. (1987) *Soil Organic Phosphorus – a Review of World Literature*. CAB International, Wallingford, 257 pp.
- Hedley, M.J., Stewart, J.W.B. & Chauhan, B.S. (1982) Changes in inorganic and organic soil phosphorus fractions by cultivation practices and by laboratory incubations. *Soil Science Society of America Journal*, 46, 970–976.
- Hieltjes, A.H.M. & Lijklema, L. (1980) Fractionation of inorganic phosphates in calcareous sediments. *Journal of Environmental Quality*, 9, 405–407.
- Hupfer, M. (1995) Bindungsformen und Mobilität des Phosphors in Gewässersedimenten (Phosphorus binding forms and mobility in lake sediments). In: Steinberg, C., Bernhardt, H. & Klapper, H. (eds.) *Handbuch angewandte Limnologie (Handbook of Applied Limnology)*, ecomed-Verlag, Landsberg am Lech, IV-3.2, 3–22 (in German).
- Jensen, H.S. & Thamdrup, B. (1993) Iron-bound phosphorus in marine sediments as measured by bicarbonate-dithionite extraction. *Hydrobiologia*, 253, 47–59.
- Joosten, H. & Clarke, D. (2002) *Wise Use of Mires and Peatlands – background and principles including a framework for decision-making*. International Mire Conservation Group and International Peat Society, 304 pp.
- Koppisch, D. (2001) Torfbildung (Peat formation). In: Succow, M. & Joosten, H. (eds.) *Landschaftsökologische Moorkunde (Landscape Ecology of Peatlands)*, Second Edition, Schweizerbart, Stuttgart, 8–17 (in German).
- Lehrkamp, H. (1989) Durch landwirtschaftliche Nutzung verursachte Veränderungen im Niedermoor, dargestellt am Beispiel des Randow-Welse-Bruches (Changes in minerotrophic peat soils due to agriculture). *Wissenschaftliche Zeitschrift der Humboldt-Universität zu Berlin, Reihe Agrarwissenschaften*, 38, 1, 12–15 (in German).
- Lewandowski, J. (2002) Untersuchungen zum Einfluß seeinterner Verfahren auf die Phosphor-Diagenese in Sedimenten (Investigations of the influence of in-lake measures on the diagenesis of phosphorus in lake sediments). Dissertation, Humboldt-Universität zu Berlin. <http://dochostrz.hu-berlin.de/dissertationen/lewandowski-joerg-2002-12-06/PDF/Lewandowski.pdf>.
- Martin, J.M., Nirel, P. & Thomas, A.J. (1987) Sequential extraction techniques: promises and problems. *Marine Chemistry*, 22, 313–341.
- Murphy, J. & Riley, J.P. (1962) A modified single solution method for the determination of phosphate in natural waters. *Analytica Chimica Acta*, 27, 31–36.
- Pagel, H., Enzmann, J. & Mutscher, H. (1982) Pflanzennährstoffe in tropischen Böden - ihre Bestimmung und Bewertung (Plant nutrients in tropical soils - determination and evaluation). VEB Deutscher Landwirtschaftsverlag, Berlin, 272 pp. (in German).
- Paludan, C. (1995) *Fosfordynamik i sedimenter fra vådområder (Phosphorus Dynamics in Wetland Sediments)*. Licentiatfhandling, Biologisk Institut, Aarhus Universitet og Afdeling for Ferskvandsøkologi, Danmarks Miljøundersøgelser, 106 pp. (in Danish).
- Parent, L.-E. & Ilnicki, P. (2003) *Organic Soils and Peat Materials for Sustainable Agriculture*. CRC Press, Boca Raton, 205 pp.
- Pettersson, K. (1986) The fractional composition of phosphorus in lake sediments of different characteristics. In: Sly, P.G. (ed.) *Sediments and Water Interactions*, Springer, Berlin, 149–155.
- Pettersson, K., Boström, B. & Jacobsen, O.-S. (1988) Phosphorus in sediments – speciation and analysis. *Hydrobiologia*, 170, 91–101.
- Psenner, R., Boström, B., Dinka, M., Pettersson, K., Pucsko, R. & Sager, M. (1988) Fractionation of phosphorus in suspended matter and sediment. *Archiv für Hydrobiologie*, 30, 98–103.
- Psenner, R., Pucsko, R. & Sager, M. (1984) Die Fraktionierung organischer und anorganischer Phosphorverbindungen von Sedimenten - Versuch einer Definition ökologisch wichtiger Fraktionen (The fractionation of organic and inorganic phosphorus compounds in lake sediments - an attempt to characterise ecologically important fractions). *Archiv für Hydrobiologie*, Suppl. 70, 111–155 (in German).

- Rump, H.-H. (1998) *Laborhandbuch für die Untersuchung von Wasser, Abwasser und Boden (Laboratory Handbook for the Analysis of Water, Wastewater and Soil)*. Wiley-VCH, Weinheim, 232 pp. (in German).
- Rupp, H., Meissner, R. & Leinweber, P. (2004) Effects of extensive land use and re-wetting on diffuse phosphorus pollution in fen areas - results from a case study in the Drömling catchment, Germany. *Journal of Plant Nutrition and Soil Science*, 167, 408–416.
- Scheffer, F. & Schachtschabel, P. (2002) *Lehrbuch der Bodenkunde (Textbook of Soil Science)*. Enke, Heidelberg, 494 pp. (in German).
- Smith, A.N. (1965) Distinction between iron and aluminium phosphate in Chang and Jackson's procedure for fractionating inorganic soil phosphorus. *Agrochimica*, 9, 162–168.
- Stevenson, F.J. (1986) *Cycles of Soil: Carbon, Nitrogen, Phosphorus, Sulfur, Micronutrients*. John Wiley & Sons, New York, 231–284.
- Succow, M. & Jeschke, L. (1990) *Moore in der Landschaft (Mires in the Landscape)*. Urania, Leipzig, 268 pp. (in German).
- Tiessen, H. & Moir, J.O. (1993) Characterization of available P by sequential extraction. In: Carter, M.R. (ed.) *Soil Sampling and Methods of Analysis*, Canadian Society of Soil Science, Lewis Publishers, Ann Arbor, MI, 75–86.
- Tischner, T. (2000) Untersuchungen zur Phosphatverlagerung und Phosphatbindung im Boden und Grundwasser einer landwirtschaftlich genutzten Fläche (Investigations on shifting and binding of phosphorus in soil and groundwater of an agricultural area). *Bodenökologie und Bodengenese*, 33, 187 pp. (in German).
- Tyler, G. (2002) Phosphorus fractions in grassland soils. *Chemosphere*, 48, 343–349.
- UNFCCC (1998) Kyoto Protocol to the United Nations Framework Convention on Climate Change. Online at: <http://unfccc.int/resource/docs/convkp/kpeng.pdf>.
- VDLUFA (1991) A 6.2.1.1: Bestimmung von Phosphor und Kalium im Calcium-Acetat-Lactat-Auszug (Phosphorus and potassium determination in the calcium-acetate-lactate digestion); A 6.2.1.2: Bestimmung von Phosphor und Kalium im Doppellactat (DL)-Auszug (Phosphorus and potassium determination in the double-lactate-digestion). In: Bassler, R. (ed.) *Methodenbuch, Band 1: Die Untersuchung von Böden*, 4, 1, Teillieferung, Darmstadt (in German).
- Velty, S., Behrendt, A. & Zeitz, J. (2006) Natural wetland restoration and the use of municipal wastewater. *Journal of Plant Nutrition and Soil Science*, 169, 642–650.
- von Post, L. (1924) Das genetische System der organogenen Bildungen Schwedens (The genetic system of the organogenic formations of Sweden). In: Comité International de Pédologie, IVème commission (commission pour la nomenclature et la classification des sols, commission pour l'Europe, président: B. Frosterus) (ed.) *Mémoires sur la Nomenclature et la Classification des Sols*, Helsingfors/Helsinki, 287–304 (in German).
- von Post, L. & Granlund, E. (1926) Södra Sveriges Torvtillgångar I (Peat resources in southern Sweden I). *Sveriges Geologiska Undersökning, Series C*, 335 (= Årsbok 19:2), Stockholm, 127 pp. (in Swedish).
- Zak, D., Gelbrecht, J. & Steinberg, C.E.W. (2004) Phosphorus retention at the redox interface of peatlands adjacent to surface waters in northeast Germany. *Biogeochemistry*, 70, 357–368.
- Zeitz, J. (2003) Bodenphysikalische Veränderungen nach intensiver Nutzung sowie nach Wiedervernässung (Soil-physical changes after intensive land use and re-wetting). In: Lenschow, U. (ed.) *Stoffausträge aus wiedervernässten Niedermooren (Discharge of substances from re-wetted minerotrophic peat soils)*. Schriftenreihe des Landesamtes für Umwelt, Naturschutz und Geologie (LUNG), Mecklenburg-Vorpommern, 1/2003, 28–37 (in German). http://www.lung.mv-regierung.de/dateien/stoffaustrag_v2.pdf.
- Zeitz, J. & Velty, S. (2002) Soil properties of drained and rewetted fen soils. *Journal of Plant Nutrition and Soil Science*, 165, 618–626.
- Zwirmann, E., Krüger, A. & Gelbrecht, J. (1999) Analytik im Zentralen Chemielabor des IGB (Analytic in the Central Analytical Laboratory of the IGB). *Berichte des IGB*, Berlin, 9, 8 (in German).

Submitted 13 Mar 2007, revision 26 Jly 2007

Editor: Olivia Bragg

Author for correspondence:

Sabine Jordan, Faculty of Agriculture and Horticulture, Institute of Crop Science, Division of Soil Science and Site Science, Humboldt-Universität zu Berlin, Invalidenstraße 42, 10115 Berlin, Germany.

Tel: +49 (0)30 2093 8622 and +49 (0)30 2093 8371; Fax: +49 (0)30 2093 8369;

E-mail: sabine.jordan@agrar.hu-berlin.de